

stacking and implies that the present tectonic structure of the E margin of the Bohemian Massif was finalized between the Upper Visean and Stephanian.

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Alpine Metamorphism in the Veporicum Unit: Differences in Reaction Mechanisms between Basement and Overlaying Sediments (Inner Western Carpathians)

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The lower metamorphic grade of the clastic Upper Paleozoic-Mesozoic cover sediments in comparison to the Alpine reworking of the basement rocks was apparent a long time ago. However, this fact did not have to necessarily reflect considerably different metamorphic condition, as suggested by Vrána (1966). Another interpretation of the different Alpine metamorphic assemblages in the basement and the cover sediments is based on a metamorphic zoning of rock pile, composed of the corresponding Veporicum, Gemicicum and ultra-Gemicicum Units (Plašienka et al. 1999). In the basement metapelites, regionally metamorphosed under prevailing amphibolite facies conditions in the Hercynian time (e.g., Zoubek, 1936), locally newly-formed garnet (enriched in grossularite component), staurolite, kyanite and tschermakitic amphibole have developed, thus reflecting peak conditions of the Alpine regional metamorphism of barrovian type (Vrána 1966, 1980; Méres and Hovorka, 1991; Kováčik et al., 1996, 1997). Absence of these higher-grade metamorphic assemblages in the Permian-Triassic clastic beds was explained by monotonous lithology of the cover rocks (Vrána, 1966). Petrographic observations on regional scale showed that the Alpine metamorphism did not establish equilibrium in the cover metasediments, as the clastic micas, plagioclase, K-feldspar are commonly preserved, and the argilliferous matrix is also frequently poorly recrystallized. Similarly, the Hercynian basement metamorphic assemblages were entirely replaced by the Alpine mineral assemblages only scarcely. Prevailing regional mineral transformations in the basement metapelites include: replacement of pre-Alpine garnet (almandine-spessartine-pyrope) by chlorite, biotite and local grossular-rich rims; chloritization or recrystallization of biotite in fine-grained biotitic mass; sericitization of plagioclase; decomposition of staurolite (or rare Al-silicates) giving rise to white micas and chloritoide. Amphibolites are retrogressed in this manner: chloritization, biotitization, epidotization, silification, albitization and the pre-Alpine Ca-amphiboles are transformed into actinolitic types.

In the overlaying Alpine units lithologic types of similar bulk-composition as within the basement metamorphites also occur. For example, metamorphic growths of white micas, chloritoide and rare kyanite (Vrána, 1964) are mainly linked to lithologies rich in pelitic compound, which occurred in the over-thrust Gemicicum Carboniferous Unit (s.l.) and sometimes in the Veporicum Permian cover-rocks. The absence of the highest grade Alpine metamorphic minerals – staurolite, garnet, (biotite) - could have been caused by a higher water content in these sediments, which were not enough dehydrated prior to the Alpine metamorphism. In the Alpine times, these rocks un-

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