

nic arc environment. Most of the classification diagrams used, are based on Rb, the element which can be very mobile, especially during metamorphism. Some varieties of the gneisses from the Lipowe Hills massif have their equivalents in the adjoining area of the Strzelin crystalline massif (SCM) (Oberc Dziedzic, 1988, 1995). Two types of the gneisses from the SCM were dated, giving the ages of 568 ± 6 Ma (the Strzelin gneiss, Oberc-Dziedzic et al., 2001) and 504 ± 3 Ma (the Gościęcice gneiss, Oliver et al., 1993). The geochemical composition of the Strzelin gneiss (Szczepański, 1999) is very similar to analysed gneisses from the northern part of the Lipowe Hills.

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Visualization of Paleostress Analysis

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In connection with work on paleostress analysis of the Roňá mineral deposit a new requirement for visualization of some result appears. Modern computers extend our possibility to present stress data in graphical form, so two new computer programs could be developed (Orient, WinOr) and direct and inverse tasks could be visualized.

If some stress tensor is known (e.g. $\sigma_1=200$ MPa in 230/30; $\sigma_2=100$ MPa in 123/26 and $\sigma_3=10$ MPa in 0/48, see Fig. 1a), it is easy to compute value of total, normal and shear stresses on any oriented plane (see Fig. 1b, 1c and 1d). Using value of internal friction and shear strength (e.g. $\phi=30^\circ$, $\tau_0=15$ MPa), we are able to evaluate differences between shear stress and critical shear stress for any plane (Fig. 1e). The zero line (wide black line in white zone) divides considered equal area plot into two parts with either positive (inside) or negative (outside) deviation. The first one is a field with reactivated planes and the second one defines a field of stability. Very illustrative is equal area projection of pitch angle. Distributional function of this angle between a strike line of a fault plane and resulted slip

vector has two extreme – 0° and 90° . The first one (zone of strike-slips, black zone in Fig. 1f) divides a plot into fields of normal and inverse faults and the second one (zone of dip-slips, white zone in Fig. 1f) into fields of sinistral and dextral faults.

Much more interesting is graphical presentation of paleostress inverse problems. In this case, we know fault plane, slip vector and sense of movement (e.g. plane 45/45, striae 45/45, thrust; see Fig. 2a). Based on some complicated mathematic analysis, it was possible to derive special equation for normalized shear stress $\tau' = \tau/\tau_{\max}$ in dependence on Lode parameter $m = (2\sigma_2 - \sigma_1 - \sigma_3)/(\sigma_1 - \sigma_3)$ and direction of main normal stresses in fault coordinate system lmn (l – parallel to striae, n – perpendicular to fault plane). In equal area projection we can plot distribution for τ' for any direction of σ_1 , σ_2 and σ_3 . The considered function gives two solutions for any s1-direction (for $m = 0.5$ see Fig. 2b and 2c) and similarly two solutions for any s3-direction (for $m=0.5$ see Fig. 2e and 2f). Very attractive is equal area plot with distribution of τ' in relationship to s2-direction. Derived function leads to one solution (for $m = 0.5$ see Fig.

2d), but the function has one point of singularity. This point is equivalent to m-direction (which is perpendicular to M-plane) and is very important for estimation whether some inverse

solution is possible or not. Geologists usually prefer s1 and s3 in their speculations, but if we understand stress problems

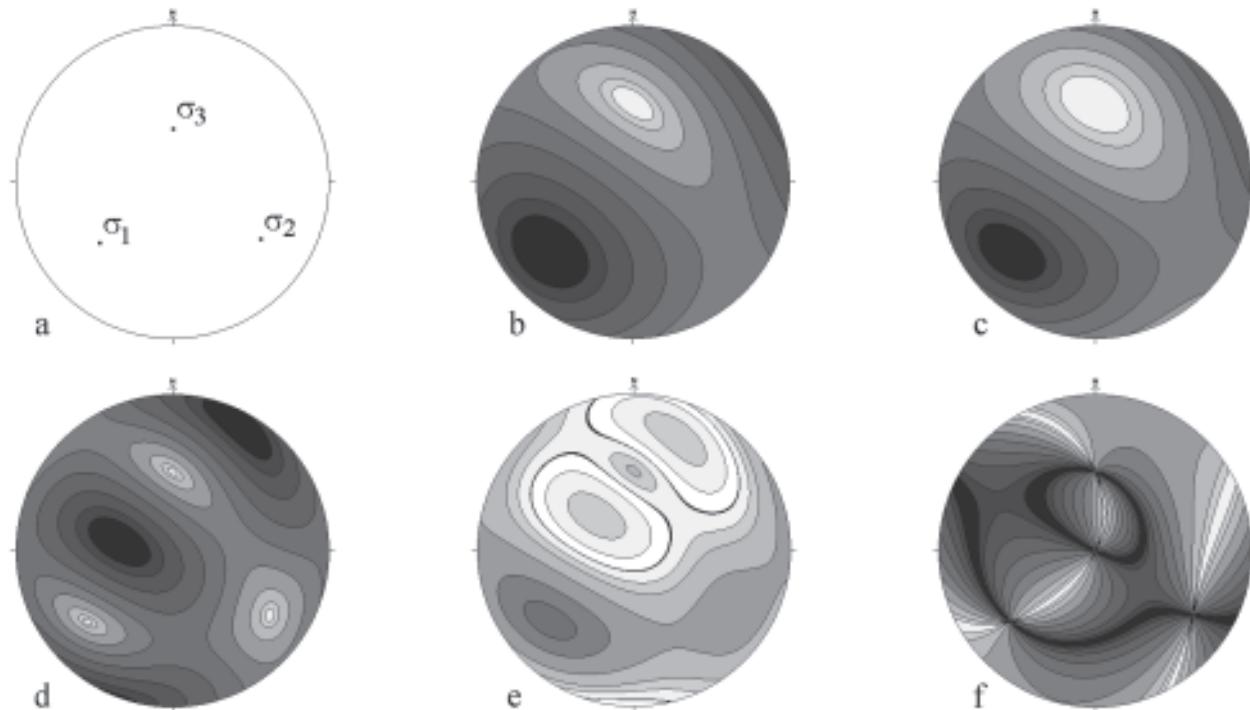


Fig. 1. Equal area projection with distribution of total stress (b), normal stress (c), shear stress (d), critical shear stress (e) and angle between slip vector and strike line (f) for one orientation of stress tensor (a). See text for explanation.

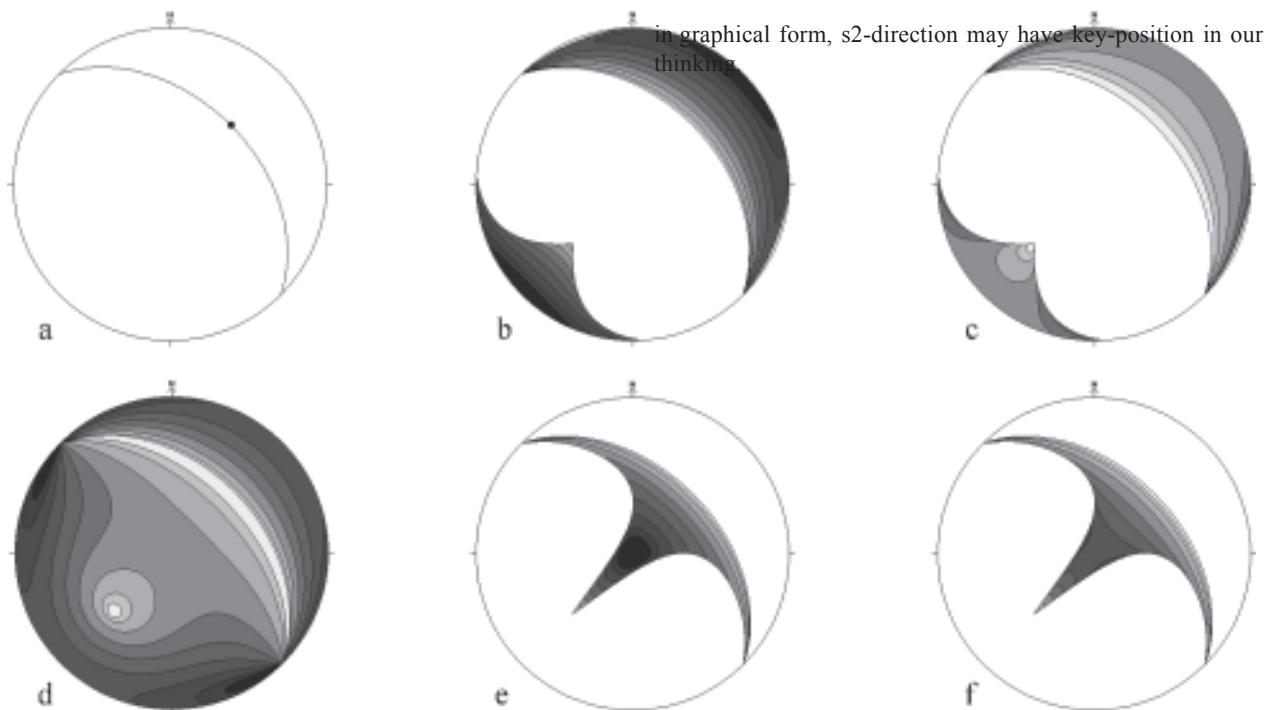


Fig. 2. Equal area plots with distribution of normalized shear stress $\tau' = \tau/\tau_{\max}$ on one fault plane (a) for different directions of s1 (b, c), s2 (d) and s3 (e, f). See text for explanation.