References

- ALIBERT C., LETERRIER J., PANASIUK M. and ZIM-MERMANN J.L., 1987. Trace and isotope geochemistry of the alkaline Tertiary volcanism in southwestern Poland. *Lithos*, 20: 311-332.
- BIRKENMAJER K., 1986. Stages of structural evolution of the Pieniny Klippen Belt, Carpathians. *Studia Geol. Pol.*, 88, 7: 32.
- BLUSZTAJN J. and HART S.R., 1989. Sr, Nd, and Pb isotopic character of Tertiary basalts from southwest Poland. *Geochim. Cosmochim. Acta*, 53: 2689-2696.
- GOLONKA J., OSZCZYPKO N. and ŚLĄCZKA A., 2000. Late Carboniferous – Neogene geodynamic evolution and paleogeography of the circum-Carpathian region and adjacent areas. Ann. Soc. Geol. Pol., 70: 107-136.
- GUTERCH A. et al., 2001. New deep seismic studies of the lithosphere in Central Europe. POLONAISE'97 and CELEBRATION 2000 seismic experiments. *Biul. Państw. Inst. Geol.*, 396: 61.
- JANKOWSKI J., TARŁOWSKI Z., PRAUS O., PECOVA J. and PETER V., 1985. The results of deep geomagnetic

sounding in the west Carpatians. *Geoph. J. R.A.S.*, 80: 561-574.

- KOVÁČ M. et al., 1998. Palinspastic reconstruction of the Carpathian-Pannonian region during the Miocene. In: M. RAKÚS (Editor), Geodynamic development of the Western Carpathians. Bratislava.
- LUSTRINO M., MELLUSO L. and MORRA V., 2002. The transition from alkaline to tholeiitic magmas: a case study from the Orosei-Dorgali Pliocene volcanic district (NE Sardinia, Italy). *Lithos*, 63: 83-113
- PLAŠIENKA D., 1999. Tectonochronology and paleotectonic model of the Jurassic – Cretaceous evolution of the Central Western Carpathians. Veda, Bratislava.
- WILSON M. and PATTERSON R., 2001. Intraplate magmatism related to short-wavelength convective instabilities in the upper mantle: evidence from the Tertiary-Quaternary volcanic province of Western and Central Europe. In: R.E. ERNST and K.L. BUCHAN (Editors), Mantle plumes: their identification through time. *Geol. Soc. Am. Spec. Paper*, 352: 37-58.

Magnetic Fabric of the Veľká Fatra Mts. Granitic Pluton – A Pilot AMS Study

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The study of the anisotropy of magnetic susceptibility (AMS) of rocks provides a useful tool for structural geology and tectonics. In the Central Western Carpathians (CWC), basic results from magnetic fabric studies of most granitoid core mountains are available (for summary see Hrouda et al., 2002). However, AMS data from the Lubochňa granite pluton of the Veľká Fatra Mts. are still missing. The Veľká Fatra Mountains typify the Core Mountains of the Tatricum, a major tectonic unit in the CWC. Their crystalline basement is represented by the Ľubochňa granitoid massif, consisting of four principal Hercynian granitic rock types building a multistage composite massif: the Smrekovica tonalite (ST) with xenoliths and wall rocks of paragneisses and orthogneisses, the Kornietov granodiorite (KGD), the Lipová granite (LG) and the Ľubochňa leucogranite (LLG). Field study as well as petrological and geochemical investigations revealed relative independence of the above granite types (Kohút, 1992) that reflect differences in the evolution of the Hercynian orogeny in the study area. Lower Carboniferous ages of magmatism of 340 ± 2 Ma were found for the KGD and 356 ± 25 Ma for the LG (Kohút et al., 1997) or 337 ± 9 Ma using cathodoluminescence-controlled single-grain (CLC) method by TIMS, as well as Ion-Microprobe (Poller et al., 2000). On the contrary, Upper Carboniferous age of 304

 ± 2 Ma was detected for the ST (Poller et al., l.c.). In addition, magnetic properties (susceptibility and NRMP) differing in individual types of granitic rocks were revealed by Kohút (1992).

In this contribution we present results of our recent AMS investigation of 251 orientated samples from 23 localities of the Veľká Fatra Mts. The measurements were performed on the KLY-3S Kappabridge (Jelínek and Pokorný, 1997). Bulk magnetic susceptibility of the ST (84 samples/8 localities) is the highest among all the rocks investigated. It ranges from 326×10^{-6} to $5,270 \times 10^{-6}$ [SI] with the mean value being $2,200 \times 10^{-6}$ [SI], suggesting I-type character of this granitoid. Bulk magnetic susceptibility of the KGD (67 samples/6 localities) ranges from 46 to $3,960 \times 10^{-6}$ [SI] with the mean value being 620×10^{-6} [SI] and lying within the susceptibility range of S-type granites. Bulk susceptibility of the LG (90 samples/8 localities) is slightly lower, ranging from 23 to $1,570 \times 10^{-6}$ [SI] with the main value being 160×10^{-6} [SI].

The degrees of magnetic lineation as well as magnetic foliation are relatively low in most of the samples, ranging from 1.02 to 1.05, with the foliation degree being slightly higher (1.05-1.10) only in ST. Planar magnetic fabric prevails in all granitoid types. In the ST, the magnetic lineation forms a welldefined maximum with the mean direction oriented WSW– ENE, plunging 10° to ENE. Also the magnetic foliation poles are well concentrated with a maximum indicating NW–SE direction, and flat plunges of 10° to 20° to the NE. In the KGD, magnetic lineation forms an irregular girdle orientated NE–SW and plunging SE at moderate angles. The magnetic foliation poles form an imperfect girdle with three conjugated submaxima orientated in a NW–SE direction. In the LG, magnetic lineation forms a well-constrained irregular girdle orientated in the NE–SW direction, plunging to the NW at a low angle. The magnetic foliation poles form a sharp anisometric maximum with a tendency to incipient steep girdle directed NW–SE.

Our AMS investigations revealed slightly different magnetic fabrics in individual granite types supporting an idea of multistage pulse character of the L'ubochňa massif of the Veľká Fatra Mts. inferred from field, petrological and geochemical studies. These differences also testify that the Alpine overprint of the magnetic fabric of the Hercynian granite frequent in other parts of the Western Carpathians (Hrouda et al., 2002) was only weak, if any, in the Veľká Fatra Mts. This enables to identify the original Hercynian magmatic fabric to a large extent. The ST is interpreted as an Upper Carboniferous metaluminous intrusion of the I-type granite into the Lower Carboniferous peraluminous S-type granite of the crystalline basement.

References

- HROUDA F., PLAŠIENKA D. and GREGOROVÁ D., 2002. Assumed Neogene deformation in the Central Western Carpathians as inferred from magnetic anisotropy investigations. EGS Stephan Mueller Spec. Publ. Series, Vol. 1: 195-211.
- JELÍNEK V. and POKORNÝ J., 1997. Some new concepts in technology of transformer bridges for measuring susceptibility anisotropy of rocks. *Phys. Chem. Earth*, 22: 179-181.
- KOHÚT M., 1992. The Veľká Fatra granitoid pluton an example of a Variscan zoned body in the Western Carpathians. In: J. VOZÁR (Editors), The Paleozoic geodynamic domains of the Western Carpathians, Eastern Alps & Dinarides. GÚDŠ Bratislava, Spec. Vol. IGCP-276, pp. 79-92.
- KOHÚT M., TODT W., JANÁK M. and POLLER U., 1997. Thermochronometry of the Variscan basement exhumation in the Veľká Fatra Mts. (Western Carpathians, Slovakia). EUG 9, Strasbourg, J. Conf. Abs., 2, 1, p. 494.
- POLLER U., TODT W., KOHÚT M. and JANÁK M., 2000. Geochemical and isotopic characteristics of the granitoids from the Velka Fatra (W-Carpathians) in combination with U-Pb Ion-Microprobe and TIMS Single Zircon Dating. Goldschmidt 2000, Oxford, J. Conf. Abs., 5, 2, p. 811.

Upper Morava and Nysa Pull-apart Grabens: Implication for Neotectonic Dextral Transtension on Sudetic Faults System

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Pull-apart basins are an integral part of intraplate and interplate strike-slip fault zones (Sylvester, 1988). Bends or sidesteps (jogs) in the main strike-slip fault system generally produce zones of extension (pull-apart basins) at releasing bends or sidesteps. Releasing sidesteps or jogs characteristically produce rhombic pull-apart basins in the overlying sedimentary section (e.g., Aydin and Nur, 1985; Sylvester, 1988). Published synoptic models usually depict pull-apart basins as simple rhombohedra bounded by two vertical, laterally offset strike-slip faults linked by two steep, parallel, oblique-slip extensional faults (e.g., Crowell, 1974; Aydin and Nur, 1985). The experiments made by Dooley and McClay (1997), however, show that pullapart basins are significantly more structurally complex than these previous synoptic models.

The performed morphotectonic analysis of the Upper Morava and Nysa pull-apart grabens and adjacent areas is based on 3D visualization and interpretation of digital elevation models (DEM), which have been compiled on the basis of detailed digitization of topographic map of 1:50,000 scale. DEM of particular areas were integrated using general DEM of the present relief of the Bohemian Massif based on GTOPO30 database (http://edcdaac.usgs.gov/gtopo30/gtopo30.html). The interpretation of DEM was compared with field mesoscopic structure-tectonic and paleostress analysis of fault populations. Special attention was focused on the architecture and kinematics of the main Sudetic fault system in the studied and adjacent areas.

The genesis and tectonic activity of the Nysa and Upper Morava pull-apart grabens are genetically related to dextral strike-slip kinematic on WNW–ESE- and NW–SE- striking fault zones, namely the Intra-Sudetic Fault, Bušín Fault, Temenice Fault and Konice Fault. Long-term dextral wrench movements and/or dextral tangential (transtensional) stress fields on the above mentioned fault systems has been sufficiently documented from the Late Variscan tectogeny to the neotectonic and Recent stages (e.g., Grygar et al., 1993; Aleksandrowski et al., 1997; Uličný, 2001; Kontny, 2001, etc.). A typical feature of the Sudetic fault system is the *en echelon* pattern of fault network. Releasing sidesteps or jogs superimposed on individual