

Magnetic susceptibility of the West Carpathian granites is in general low, in the order of 10^{-4} [SI]. In the minority of specimens it is in the order of 10^{-5} and in exceptional specimens it is higher, in the order of 10^{-3} . The susceptibility values of the most West Carpathian granites correspond to the values

typical of S-types. This is in contradiction with the granite origin revealed geochemically. The preliminary explanation of this contradiction is that magnetite originally present in the I-types was destroyed during Alpine deformation indicated by magnetic fabric.

Development of the Gföhl Migmatites through Partial Melting and Textural Annealing of High-Grade Orthogneiss via Process of Disintegration of Solid State Texture

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The Gföhl gneiss is an important lithological unit of the most deeply buried part of the Moldanubian root domain. It is suggested by some authors that this gneiss complex together with HP granulites, eclogites and mantle fragments form the highest structural position of the Moldanubian zone due to exhumation associated with large-scale nappe tectonics (Matte, 1990, Petrakakis, 1995). The Gföhl gneiss complex is rather heterogeneous in composition, and consists of nebulitic migmatites, stromatitic biotite-rich migmatites and locally of banded orthogneisses. The detailed outcrop observations reveal gradual transitions from porphyritic orthogneiss via high-grade mylonites to entirely molten rock. The question arises, whether the large volumes of felsic medium-grained migmatites originate through partial melting of orthogneiss protolith or whether the main proportion of migmatitic domain was originally formed by fertile metasediments rich in hydrous minerals (Thompson, 2000).

Several key outcrops were investigated from the structural, textural and petrological point of view. The earliest structures are represented by steep solid state high-grade foliations in strongly sheared orthogneisses marked by alternations of monomineralic feldspar and quartz layers and large amount of biotite. This first fabric is W dipping (at an angle of 60–80°) and N-S trending. The early fabric is folded and transposed by flat foliation associated with the development of fine-grained mylonites. These high-grade mylonites progressively pass into migmatitic gneiss rich in sillimanite and garnet, and finally into completely molten felsic leucosomes. The new planar system is E-W trending (at an angle of 15–40°) and bears N–S oriented mostly subhorizontal lineation. We sampled two sections in which the above transition from banded orthogneiss to nebulitic migmatite was observed. The aim of this work is to understand the melting process of high-grade orthogneiss, i.e., relatively refractory rock using a detailed petrological and textural analysis.

We have distinguished and documented four textural stages from the orthogneiss to nebulitic migmatites. The first one is represented by fine-grained banded orthogneiss with distinctly separated monomineral layers. The K-feldspar layers

0.75–2 mm thick consist of grains 0.5 mm in size with straight boundaries. Numerous rounded inclusions of quartz (0.01 mm in size) occur mostly at triple points. A polygonal mosaic of well-equilibrated plagioclase 0.2–0.3 mm large forms layers 0.25–1.25 mm thick. Quartz occurs in 0.7–0.1 mm thick polycrystalline ribbons. Large flakes of biotite, locally overgrown by sillimanite (< 5%), form bands separating quartz from plagioclase aggregates. Small garnet (0.07–0.1 mm in size) is associated with biotite aggregates.

The second stage is characteristic of grain coarsening and disappearance of monomineralic layering. The K-feldspar-rich aggregates are composed of K-feldspar (80%) grains (0.6 mm in size) with straight boundaries and numerous inclusions of quartz (20%) and biotite. Plagioclase-rich layers are composed of plagioclase (80%), quartz (20%), biotite flakes, and rounded garnet grains (0.5–0.9 mm in size) + sillimanite. Quartz forms irregular aggregates composed of large grains (0.4–0.6 mm in size) with strongly lobate boundaries.

The third stage is marked by change in the proportion and size of individual minerals and by the increase in sillimanite content. Former plagioclase-rich layers show almost granite-like texture being composed of almost equivalent amount of plagioclase and quartz as well as minor amount of K-feldspar. The K-feldspar layers consist of large irregular grains of K-feldspar and small plagioclase grains. The K-feldspar/quartz ratio is 1:1. The plagioclase and K-feldspar layers are separated by sillimanite aggregates. No relics of original layering can be observed in the nebulitic migmatite.

The grain boundaries of minerals in individual stages were “traced” in the GIS Arc View environment and analysed using The Matlab™ Poly LX toolbox (Lexa, 2001), where statistical analysis of grain size, grain contact frequencies, modal compositions and grain boundaries and shapes were performed. This study permits to quantify the textural evolution of mineral aggregates from solid state to random anatectic structure. The detailed microprobe work, currently in progress, is carried out to estimate the PT conditions of textural annealing and melting. Basing on these data, an attempt is made to propose a model of textural annealing and melting explaining the disintegration of

the original solid state texture and development of large volumes of nebulitic migmatites in the Gföhl unit.

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Geochemical Variability of the Kłodzko – Złoty Stok Massif: Possible Role of Multiple Mafic End-Members of Hybrid Granitoids

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The Kłodzko - Złoty Stok massif (KZS) is one of the Variscan plutonic bodies located within the Lugian Zone at NE margin of the Bohemian Massif in southern Poland. Though the rocks were described petrographically in a great detail (namely by Wierzcholowski 1976) and many silicate analyses were published, the trace element contents and geochemical characteristics are still poorly known.

The plutonic rocks from the KZS are dominantly intermediate and only marginally acidic. The whole compositional range of major plutonic rocks (excluding enclaves) is from about 54 to 68 wt.% with rare leucogranites up to 75%.

Lorenc (1991) stressed the metaluminous chemistry of the major rock types, their hybrid nature and abundance of dark inclusions that he interpreted as typical mafic magmatic enclaves (MME). According to him the role of mafic magma in petrogenesis of granitoids forming the KZS was principal.

Geochemical study of a new set of samples displays highly variable composition of major rock types as well as of MME. These data enabled us to differentiate at least 3 compositional groups (Table 1) that differ in petrochemical parameters.

Mafic varieties of rocks from the KZS display typical features of mantle-derived magmas or hybrid magmas dominated in composition by the mantle end-members (high mg-values, high contents of MgO, Cr, Ni). However, composition of the

mafic members vary and they cannot represent single magma batch.

The most potassic plutonic rock yet analysed are monzonites from the endocontact of the KZS at Żelazno. These and some other dark and K-rich rock varieties from KZS are similar in chemical composition to “vaugneritic” and “syenitic” intrusions in the Niemcza Zone, namely at Koźmice and Piława Górna (cf. Puziewicz 1987, 1988). Their chemical composition cannot be due to contamination of a common basaltic magma with crustal rocks or melts as these rocks have not only high K and Rb but also the highest MgO, Cr and Ni and the mg-values.

Composition of the prevailing relatively dark and K-rich granitoids resemble that of shoshonitic rocks (SHO). Some granitoids from surroundings of Laskówka and Laski correspond to the high-K calc-alkaline series (HKCA) with different composition of MME.

Compared to durbachitic rocks from the Moldanubian Zone of the Bohemian Massif, even the most potassic rocks from the KZS and the Niemcza Zone display significantly lower contents of K₂O, P₂O₅, Rb, Cs, Th and U. However, their geochemical signature, namely in the “spider diagrams”, seems to be similar. We consider existence of some similarities in history of their mantle sources.

Group	Rocks	SiO ₂	K ₂ O	K ₂ O/Na ₂ O	mg
UK	monzonite to melagranodiorite	54–58	3.4–4.7	1.3–2.2	72–65.3
SHO	melagranodiorite to granodiorite	58.7–62.5	3.4–4.2	1.1–1.3	55.7–53.1
HKCA	biotite granodiorite to monzogranite	64–68	3.1–3.5	1.0–1.1	46.4–43.1

mg = MgO/(MgO + FeOtot.) (from molar values)

Tab. 1. Comparison of selected petrochemical parameters for major compositional groups of plutonic rocks from the Kłodzko – Złoty Stok massif and similar rocks from the Niemcza Zone.