Alpine-type deformations with isoclinal and nappe structures (e.g. thrust of the presumed Ordovician complex over the Devonian-Early Carboniferous one in the central part of the Ještěd Mts., as interpreted already by Gallwitz, 1930) seems to be most probable.

### References

CHLUPÁČ I. 1993. Stratigraphic evaluation of some metamorphic units in the N part of the Bohemian Massif. Neues Jahrbuch für Geologie und Paläontologie, Abhandlungen, 188, 363-388.

 1997. Palaeozoic ichnofossils in phyllites near Železný Brod, northern Bohemia. Journal of the Czech Geological Socie-

ty, 42, 75-94.

- (submitted). Poznámky k rozšíření devonu a stavbě metamorfovaného paleozoika v jižní a střední části Ještědského pohoří. Zprávy o geologických výzkumech v roce 1997.
- CHLUPÁČ I. and HLADIL J. 1992. New Devonian occurrences in the Ještěd Mts., North Bohemia. Časopis pro mineralogii a geologii, 37, 185-191.
- GALLWITZ H. 1930. Geologie des Jeschkengebirges in Nordböhmen. Abhandlungen des Sächsischen Geologischen Landesamts, 10, 1-63.
- GUNIA T. 1967. Cambrotrypa (Tabulata) z metamorfiku Sudetów Zachodnich. Rocznik Polskiego towarzystwa geologicznego, 37, 417-428.

# Different Geodynamic Setting of Subunits in the HP/LT Metamorphosed Bôrka Nappe (Inner Western Carpathians): Evidence from Basic Rocks Geochemistry and Petrology

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The newly defined Bôrka Nappe (Mello et al. 1990) is located in the northern part of Slovenský kras Mts. and in the western part of the Spišsko-gemerské rudohorie Mts. It represents a HP/LT metamorphosed part of an ancient accretionary prism formed by closing the Meliata-Hallstatt ocean during the Jurassic time. The Bôrka Nappe is recently stratigraphically divided into three formations: (1) Jasov and (2) Bučina Formations, probably Permian in age, composed of metamorphosed clastic sediments and acid volcanoclastics, and (3) Dúbrava Formation, supposedly Mesozoic in age, containing metamorphosed sedimentary and basic magmatic rocks. Relic primary textures and geochemical and petrological characteristics of these basic magmatic rocks have been studied. Obtained data indicate that following five group of HP/LT metamorphosed basic rocks, which differ in protolith, geochemical signature and metamorphic evolution, can be recognised in the Dúbrava Formation:

- (1) Basalts, dolerites and gabbros with BABB (back-arc basin basalt) to N-MORB (normal oceanic-ridge basalt) affinity transformed into epidote-glaucophane rocks with variable amounts of Na-pyroxene (acmite), clinozoisite, albite, white mica and garnet. Basalts, originally forming lava flows with glassy or lava breccia rinds, experienced a progressive metamorphic evolution from prehnite-pumpellyite trough greenschist to epidote blueschist facies conditions. Vestiges of latter ocean-ridge type metamorphism have been found in the dolerites and gabbros.
- (2) Dolerites and less amount of basalts still preserved dolerititic or ophitic texture geochemically close to BABB, formerly metamorphosed in epidote blueschist metamorphic conditions and latter retrogressed to greenschist conditions. Actinolite, chlorite, albite and octaeder-forming magnetite is a typical mineral assemblage.
- (3) Basalts, rarely dolerites with CAB (calc-alkaline basalt) affinity. Relic porphyritic or ophititic texture is rarely preserved. The oldest metamorphic phase, probably in HP/MT conditions (high-Al amphibole relics) was overprinted by HP/LT metamorphic stage producing blueschist with vari-

- able amounts of glaucophane, garnet, omphacite, clinozoisite/epidote, white mica, chlorite and rutile. These blueschists were largely retrogressed to greenschists composed of actinolite, epidote, chlorite and albite.
- (4) Banded basaltic pyroclastics with a calc-alkaline signature retrogressed from blueschist metamorphic conditions to greenschist of actinolite-chlorite-epidote-albite composition with magnetite octaeders.
- 5) Phyllonitised (under greenschist facies conditions) amphibolites geochemically close to N-MORB with a HP/LT metamorphic overprint and mostly retrogressed again under greenschist facies conditions.

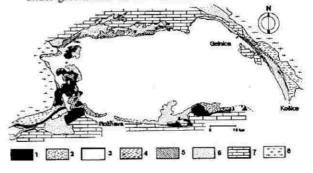


Fig. 1. 1-2 Bôrka Nappe: 1-Dúbrava Fm. 2-Jasov and Bučina Fm. 3-6: Palozoic of Gemeric Unit: 3-Early Paleozoic fomrtions 4-clastic Carboniferous formations 5-Zlatník Fm. 6-Gemeric Permian formations 7-Mesozoic formations of Silica and Stratená Nappe and Meliata Unit; black triangle - Rudník (Dúbrava Fm.)

The basic magmatic rocks of all the above-mentioned groups occur in the Bôrka Nappe separately in the association with specific types of metamorphosed sedimentary rocks. This implies that the Dúbrava Formation is, in fact, composed of several individual lithostratigraphic formations in a form of slices with mutual tectonic contacts. The supposed formations,

based on their original geodynamic setting and further tectonometamorphic evolution, belong to two essential types: (1) material of the subducted lithospheric plate of the Meliata-Hall-statt ocean and (2) material of the overriding lithospheric plate immediately above subducting oceanic plate. The first type is represented by metamorphosed basaltic rocks with BABB to N-MORB affinity (Groups 1 and 2) occurring in the association with carbonates, rarely pelitic schists and also radiolarites. All these rocks are probably Mesozoic in age. The latter type is represented by metamorphosed calc-alkaline basalts (Group 3) interlayered by pelitic metasediments containing originally variable amounts of organic matter and pyroclastic component. The age of this formation is probably Early Palaeozoic (Faryad and Henjes-Kunst 1997 - 40 Ar/39 Ar dating on

phengite). Polymetamorphosed amphibolite associated with phyllonitised gneisses (Group 5) and banded metamorphosed calc-alkaline volcanoclastics interlayered with banded carbonates (Group 4) belong probably to the same type.

#### References

MELLO J. et al. 1997. Explanation notes to the geological map of the Slovenský kras Mts. 1: 50 000. (In Slovak with English resume). Vyd. Dyonýza Štúra, 256 pp. Bratislava

FARYAD S.W. and HENJES-KUNST F. 1997. Petrological and K-Ar and <sup>40</sup>Ar-<sup>39</sup>Ar age constraints for the tectonothermal evolution of the high-pressure Meliata unit, western Carpathians (Slovakia). *Tectonophysics*, 280, 141-146.

# Sr-Nd Isotopic Composition of Granites from the Northern Part of the Moldanubian Pluton and its Significance for Genetic Classifications

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In order to characterise sources and processes that could have been involved in genesis of peraluminous granites in the northern part of the Moldanubian Pluton (MP) a combined study of whole-rock major- and trace-element geochemistry and Sr-Nd isotopic composition was undertaken. As precise geochronological information on individual rock types is still scarce, a muscovite Ar-Ar age of 328 Ma (Scharbert et al. 1997) was chosen to age-correct all isotopic ratios. The granites are classified using a scheme published by Matějka (1997 and references therein); undistinguished fine-grained granites between Čeřínek and Landštejn bodies are termed Eisgarn s.l.

The new Sr isotopic ratios (Tab. 1) tie in with those published previously (Scharbert and Veselá 1990) and thus all the available data were treated as a single data set. Taken together, they are very variable  $(^{87}Sr)^{86}Sri = 0.7049-0.7144)$ , spanning a wide compositional range from little modified, mantle or lower crustal rocks to mature metasedimentary lithologies such as Moldanubian paragneisses. The most primitive is the Pavlov granite (0.7049-0.7051, n = 2), followed by Boršov (0.7067-0.7075, n = 6), one analysis of the Světlá type (0.7075: Kam. Lhota) and Čeřínek granite (0.7090-0.7100, n = 3). Eisgarn s.l. from Horní Cerekev (Scharbert and Veselá 1990) yielded an initial ratio of 0.7108 — identical to Eisgarn granite analysed by Liew et al. 1989. Comparable composition shows majority of Lipnice type (0.7092-0.7111, n = 3) and one Světlá granite (0.7110: D. Březinka). The Sr-isotopic ratios from the Bílý kámen granite are extremely variable (0.7051-0.7120, n = 10). The most evolved strontium contains typical Eisgarn s.l. (0.7122-0.7143, n = 7).

On the other hand, the obtained initial Nd-isotopic compositions (so far not determined in this part of the MP) are all evolved and more or less uniform ( $\epsilon^i_{Nd} = -5.6$  to -7.9; Tab. 1). High two-stage Nd model ages indicate a negligible (if any) role for depleted-mantle derived magmas and a long crustal history of the protolith ( $T^{DM}_{Nd} = 1.48-1.66$  Ga). The most primitive is Lipnice (-5.6), followed by Světlá (-6.2), Pavlov (-6.5),

|      | Locality<br>(type') | 87Sr/86Sr <sub>i</sub> + | <sup>143</sup> Nd/ <sup>144</sup> Nd <sub>i</sub> * | $\epsilon_{Nd}^{i}{}^{\dagger}$ | T <sub>Nd</sub> ‡ |
|------|---------------------|--------------------------|-----------------------------------------------------|---------------------------------|-------------------|
| Ce-1 | Boršov (B)          | 0.70672                  | 0.511850                                            | -7.1                            | 1.60              |
| Ce-2 | Boršov (B)          | 0.70746                  | 0.511842                                            | -7.3                            | 1.62              |
| Ce-3 | Čeřínek (Č)         | 0.70974                  | 0.511811                                            | -7.9                            | 1.66              |
| Me-1 | D. Město<br>(L)     | 0.70921                  | 0.511928                                            | -5.6                            | 1.48              |
| Me-2 | D. Město<br>(L)     | 0.71108                  | -                                                   |                                 | <del>-</del>      |
| Me-3 | D. Březinka<br>(Sv) | 0.71099                  | 0.511896                                            | -6.2                            | 1.53              |
| Me-5 | Pavlov (P)          | 0.70493                  | 0.511883                                            | -6.5                            | 1.55              |
| Me-6 | Kam. Lhota<br>(Sv)  | 0.70754                  | :=/                                                 | *                               | 3+6               |
| Me-7 | Stvořidla<br>(St)   |                          | 0.511847                                            | -7.2                            | 1.61              |

Rock types (Matějka, 1997): B: Boršov, Č: Čeřínek, L: Lipnice, Sv: Světlá, P: Pavlov, St: Stvořidla

isotopic ratios age-corrected to 328 Ma

Tab. 1. Sr-Nd isotopic data for granites from northern MP.

Boršov (-7.1 to -7.3), Stvořidla (-7.2) and Čeřínek types (-7.9). Austrian Eisgarn has  $\varepsilon = -7.5$  (Liew et al. 1989).

The relatively primitive Sr-isotopic composition of the Pavlov and Boršov granites rules out their genesis by partial melting of ordinary Moldanubian metasediments and may point to a lower crustal source with low time-integrated Rb/Sr ratio. On the other hand, Sr-Nd isotopic signature of both the Lipnice and Světlá types, and, in particular, Eisgarn s.l., is

two-stage Nd model ages (Ga; Liew and Hofmann 1988)